

Present-day mass variations in Fennoscandia as determined from joint analysis of absolute gravity and GRACE data

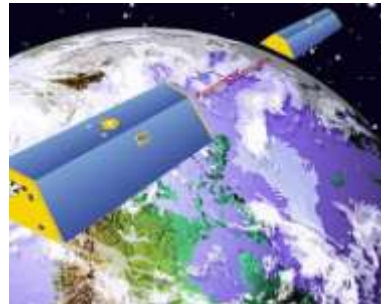
Jürgen Müller, Majid Naeimi, Olga Gitlein, Ludger Timmen and Heiner Denker



DFG Priority program SPP 1257:
Mass distribution and mass transport
in the Earth system



Centre of excellence:
Quantum Engineering and
Space Time-Research



Introduction

- GIA-induced effects are observed by geodetic techniques in Fennoscandia
- GRACE monthly solutions reflect the integral effect of mass variations in the atmosphere, hydrosphere and geosphere
- Absolute gravity and GPS campaigns have been run under the umbrella of the Nordic Geodetic Commission, also providing geometric features

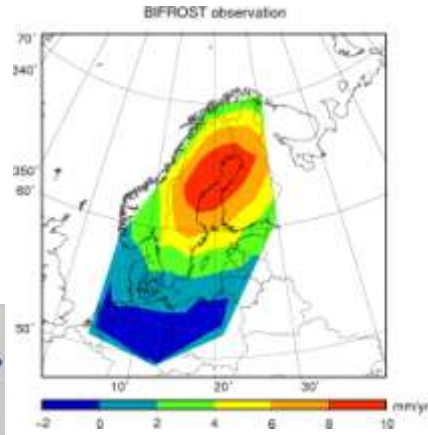
Observation of GIA-induced effects

Land uplift (1 cm / year)



Observed by

- GPS
 - GRACE
 - terrestrial gravimetry
- $\sigma_g = \pm 20 \text{ nm/s}^2$



BIFROST GPS network,
Lidberg et al. (2007)

Analysis of GRACE monthly solutions

General strategy for the computation of trends:

- Computation of grid values dg from spherical harmonic coefficients up to degree and order n
- Filtering and synthesis of a time series of grids
- Pixel-wise least-squares adjustment

$$dg(\varphi, \lambda, t) = A + Bt + \sum_{i=1}^k C_i \cos(\omega_i t) + D_i \sin(\omega_i t)$$

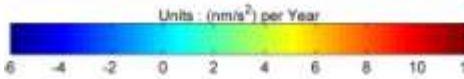
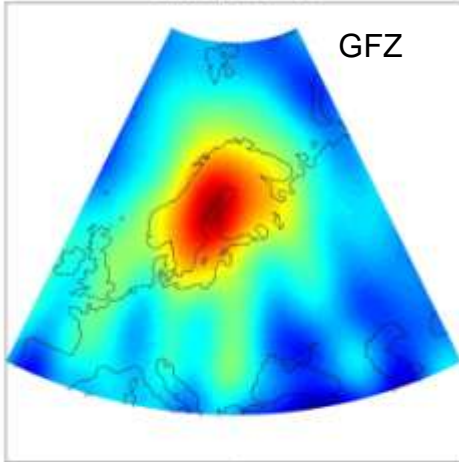
trend periodic variations

Which periods should be considered?

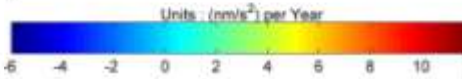
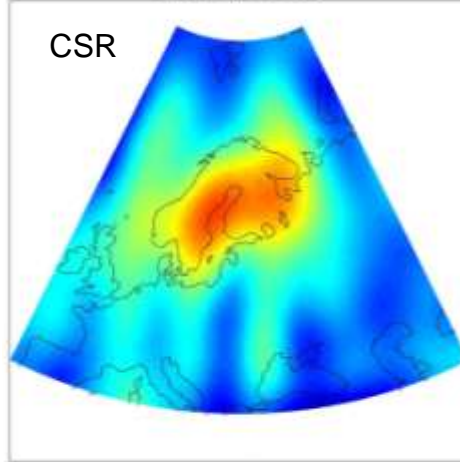
Also a quadratic term in time?

Secular trend in Fennoscandia

Secular gravity variations computed from GFZ DATA
Period : Jan 2003 - Feb 2009
Gaps: June 2003/Jan 2004
Smoothing radius = 400

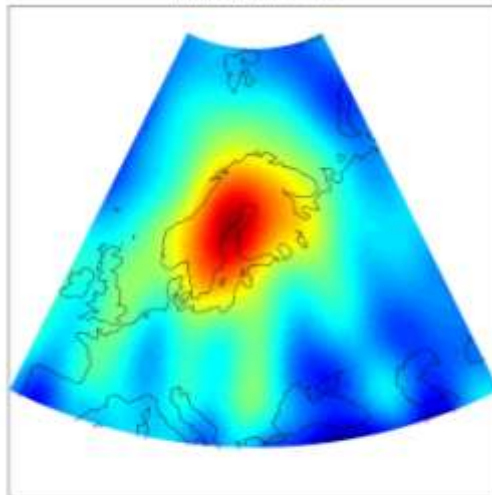


Secular gravity variations computed from CSR DATA
Period : Jan 2003 - Feb 2009
Gaps: June 2003
Smoothing radius = 400



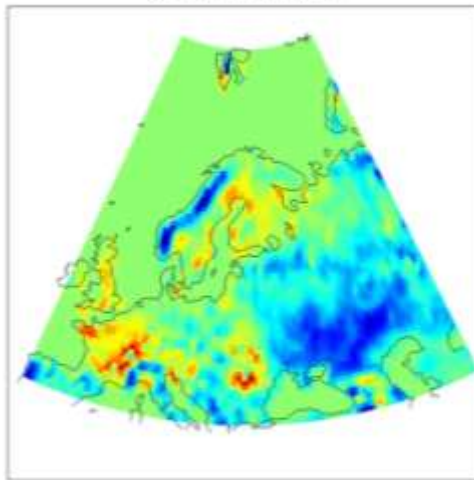
Comparison of different filter radii

Smoothing radius = 400



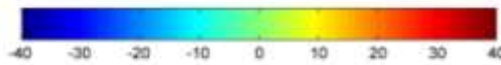
Possible hydrology contribution

Secular variations computed from GLDAS model
Soil moisture + Snow water



GLDAS,
Jan. 2003 – Feb. 2009

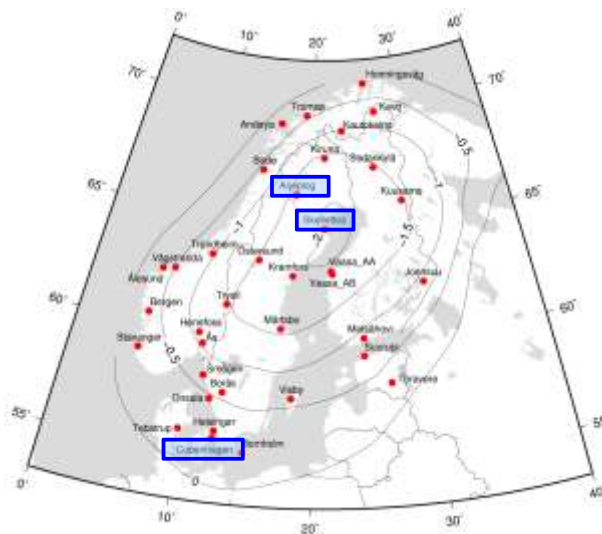
Units : mm (water) per year



GRACE gravity variations at AG sites

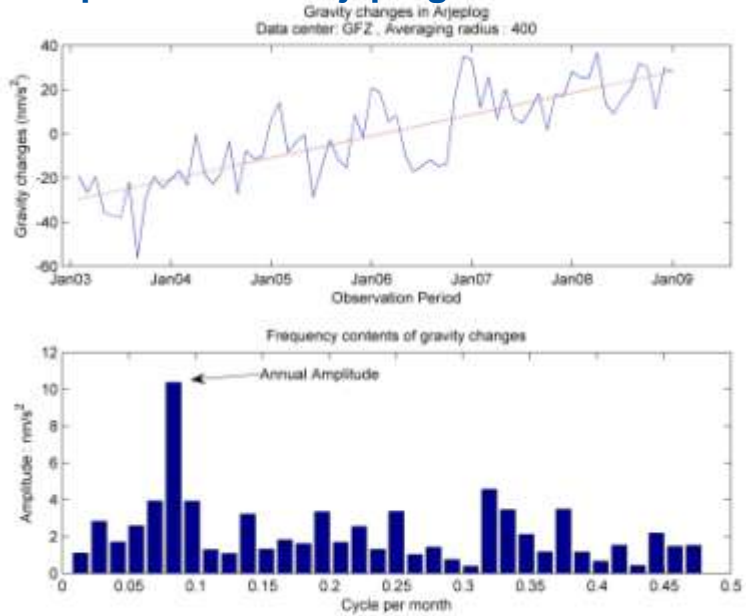


FG5-220 from IfE (Photo: Gitlein)

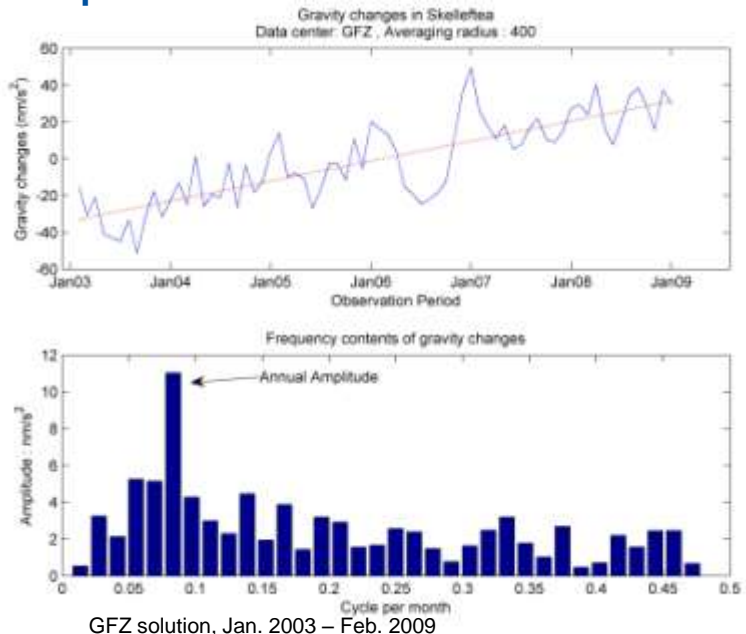


Gravity change after Ekman and Mäkinen (1996)

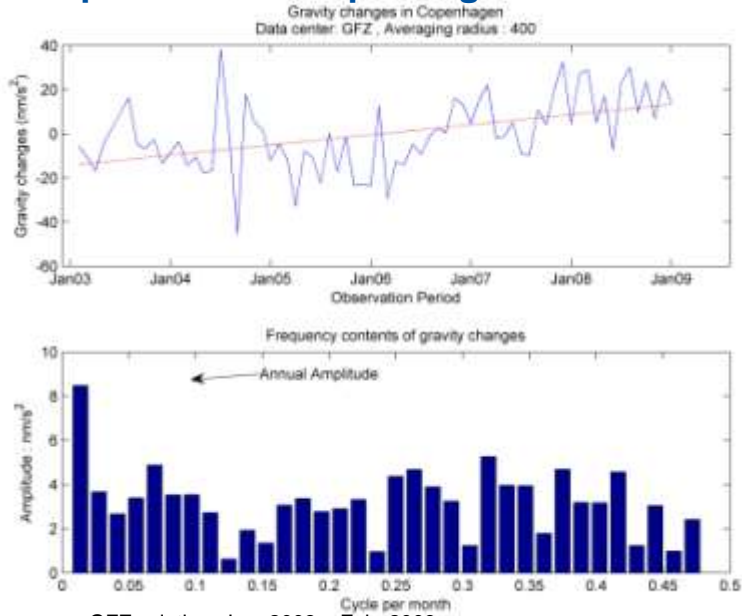
GRACE spectrum for Arjeplog



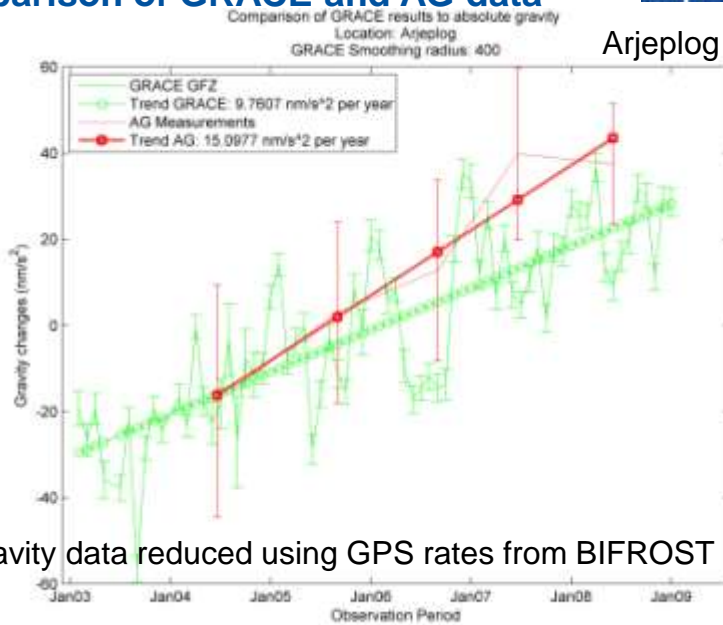
GRACE spectrum for Skelleftea



GRACE spectrum for Copenhagen



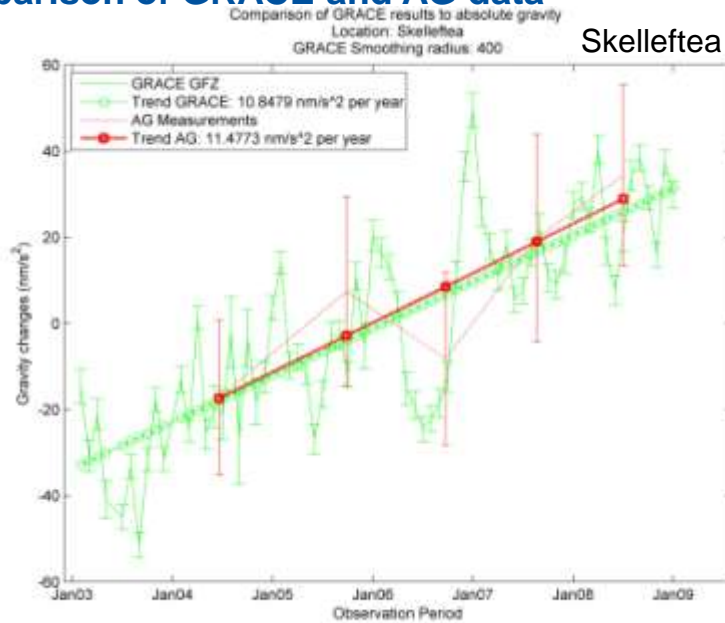
Comparison of GRACE and AG data



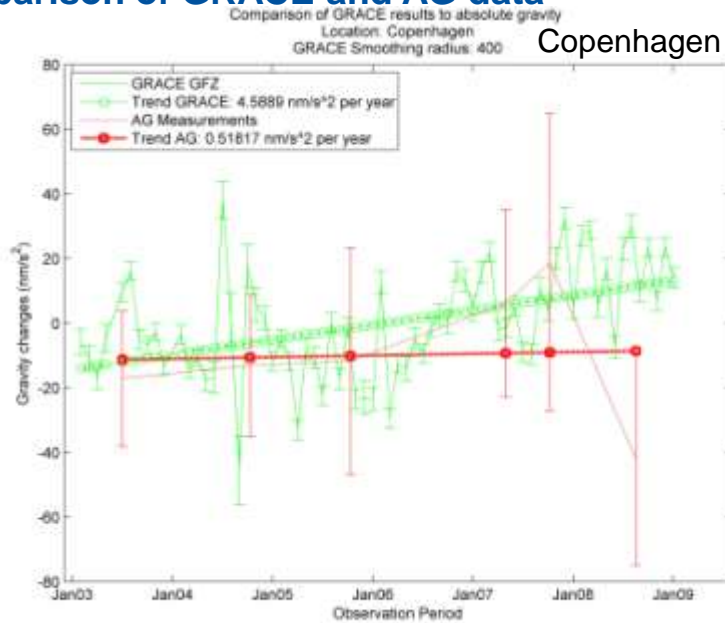
AG gravity data reduced using GPS rates from BIFROST



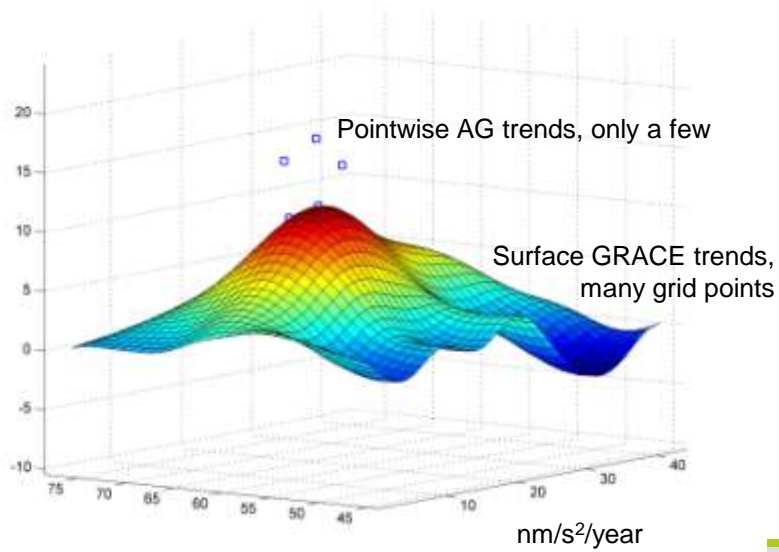
Comparison of GRACE and AG data



Comparison of GRACE and AG data

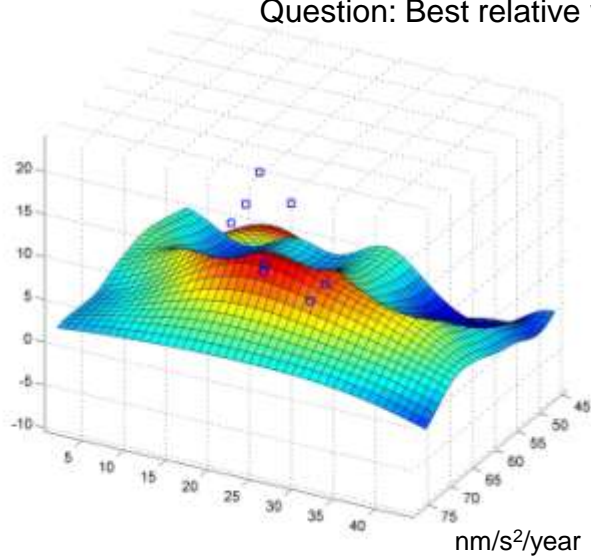


Combination of AG and GRACE, data

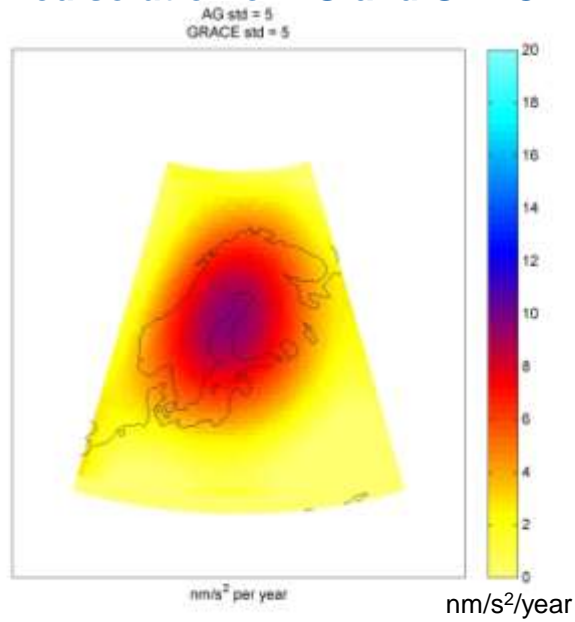


Combination of AG and GRACE, data

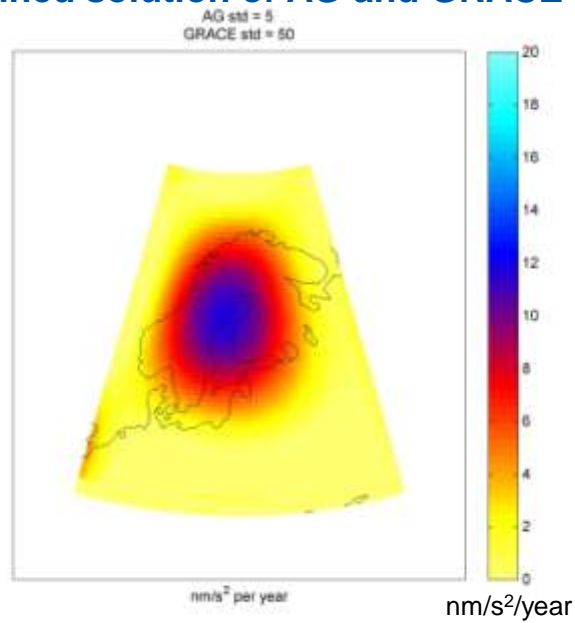
Question: Best relative weighting?



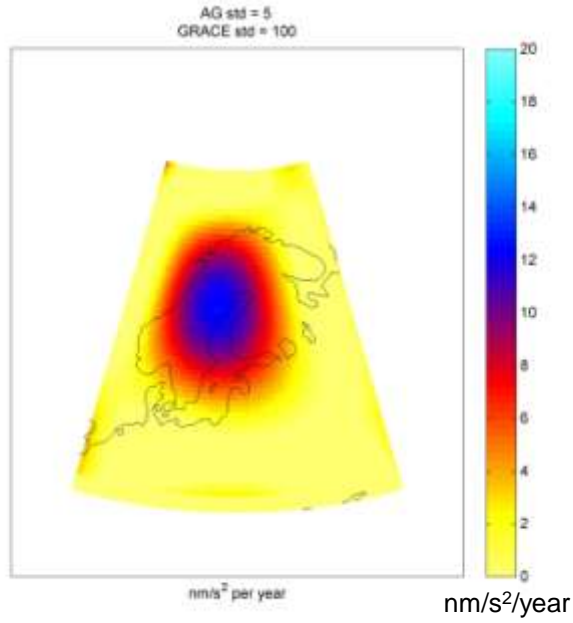
Combined solution of AG and GRACE



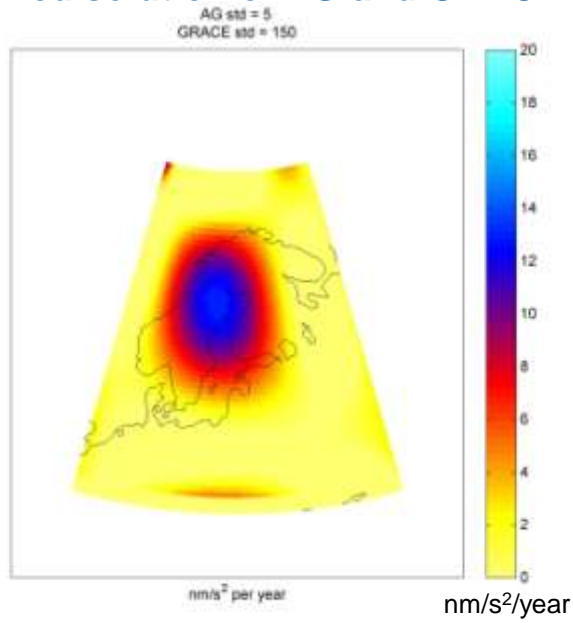
Combined solution of AG and GRACE



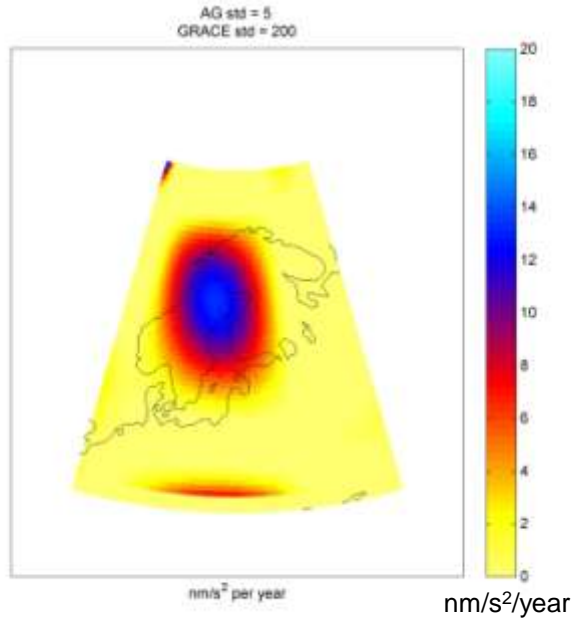
Combined solution of AG and GRACE



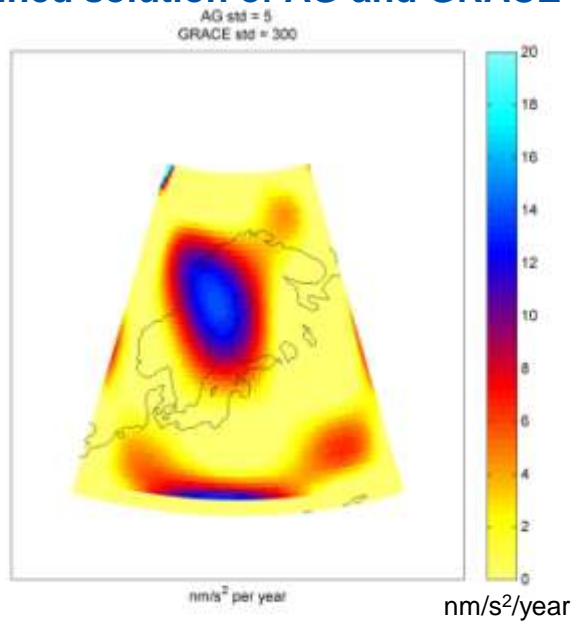
Combined solution of AG and GRACE



Combined solution of AG and GRACE



Combined solution of AG and GRACE



Conclusions

- GIA signature is significant in GRACE data, up to 11 nm/s²/year for Fennoscandia
- Uplift centre and shape comparable with terrestrial measurements such as GPS and AG
- Combined solution requires
 - more AG data
 - consistent pre-processing (hydrology)
 - relative weighting to be optimised

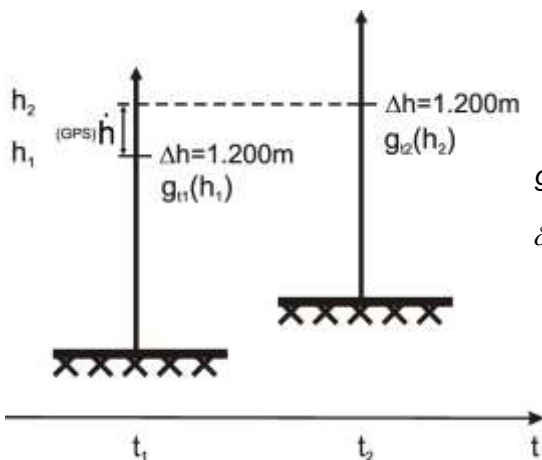
More data and refined modelling helpful!



Conversion of gravity and height changes

Geoid change:

Change of gravity disturbance:



$$\dot{N} = \frac{R}{4\pi\gamma} \iint_{\sigma} H(\psi) \left(\dot{g} + \frac{2\gamma}{r} \dot{h} \right) d\sigma$$

gravity disturbance

$$g_{11}(h_2) = g_{11}(h_1) + \frac{\partial g}{\partial h} \dot{h}$$

$$\dot{\delta g} = g_{12}(h_2) - g_{11}(h_2)$$

Example:

$$\dot{h} = 1.0 \text{ cm}$$

$$g_{12}(h_2) - g_{11}(h_1) = -2.0 \mu\text{Gal}$$

$$\frac{\partial g}{\partial h} = -3.0 \frac{\mu\text{Gal}}{\text{cm}}$$

$$\dot{\delta g} = +1.0 \mu\text{Gal} / \Delta t$$



Comparison of GRACE and AG data

Comparison of GRACE results to absolute gravity
Location: Vaasa AB
GRACE Smoothing radius: 400

Vaasa

